

SIMULATION OF ANALYTICAL TRANSIENT WAVE DUE TO DOWNWARD BOTTOM THRUST

S.S. TJANDRA^{1,2}, S.R. PUDJAPRASETYA¹, L.H. WIRYANTO¹

¹Industrial and Financial Mathematics Research Group,
Bandung Institute of Technology, Indonesia

²Industrial Engineering,
Parahyangan Catholic University, Bandung, Indonesia
ssudharmatjandra@yahoo.com

Abstract. Generation process is an important part of understanding waves, especially tsunami. Large earthquake under the sea is one major cause of tsunamis. The sea surface deforms as a response from the sea bottom motion caused by the earthquake. Analytical description of surface wave generated by bottom motion can be obtained from the linearized dispersive model. For a bottom motion in the form of a downward motion, the result is expressed in terms of an improper integral. Here, we focus on analyzing the convergence of this integral, and then the improper integral is approximated into a finite integral so that the integral can be evaluated numerically. Further, we simulate free surface elevation for three different types of bottom motions, classified as impulsive, intermediate, and slow movements. We demonstrate that the wave propagating to the right, with a depression as the leading wave, followed with subsequent wave crests. This phenomena is often observed in most tsunami events.

Key words: Analytical solution, dispersive waves, bottom motion.

Abstrak. Proses pembangkitan adalah satu bagian penting dari pemahaman gelombang, khususnya tsunami. Gempa besar di bawah laut merupakan salah satu penyebab utama tsunami. Permukaan laut terdeformasi sebagai respon dari gerakan dasar laut akibat gempa. Deskripsi analitis dari gelombang permukaan yang dihasilkan oleh gerakan dasar laut dapat diperoleh dari model linier dispersif. Untuk gerakan dasar ke bawah, formulasi analitik elevasi gelombang permukaan dinyatakan dalam bentuk integral tak wajar. Pada paper ini, kita fokus pada analisis kekonvergenan integral tersebut dan integral tak wajarnya kita hampiri dengan sebuah integral batas hingga dan selanjutnya kita hitung secara numerik. Setelah itu kita simulasikan elevasi permukaan yang timbul akibat tiga jenis gerakan dasar, yaitu gerakan impulsif, menengah, dan lambat. Hasil yang diperoleh menunjukkan bahwa gelombang yang terbentuk merambat ke kanan, diawali dengan gelombang negatif utama yang diikuti dengan deretan gelombang-gelombang kecil (ripple) di belakangnya. Fenomena semacam ini sering kita temui pada kebanyakan peristiwa tsunami.

Kata kunci: Solusi analitik, gelombang dispersif, gerakan dasar.

1. INTRODUCTION

Basically, tsunami events consist of three important stages. The stages are generation, propagation and run up. The last two stages are strongly influenced by the first stage. Hence, studying tsunami generation process is very important. Here, we will study the formation of free surface elevation due to downward bottom motion. In particular, we study the relation between bottom motion characteristics and the amplitude of resulting waves.

Study on the surface wave generated by bottom motion has long been an interesting subject of research. Hammack [1, 2] was among the first who did notable research in this area. He conducted careful experimental works on vertical bottom motions. Until recently, his experimental results has still been used as the benchmark test for numerical models, see for instance [6] and [7]. Hammack also studied and compared analytical solution with his experimental and numerical results from KdV model. Dutykh [4] and Kervella et.al. [5] studied tsunami generation using linear and non linear models.

An analytical description of transient wave as response of an exponential bottom disturbance is recorded in [1]. This exponential bottom motion can be used for simulating various types of motions, including impulsive and slow motion. In the analytical formula, the surface wave is written in terms of improper integral. But finding the anti derivative of this integral is not straightforward, therefore here we implement numerical integration. After analyzing its convergence and determine the representative integral, we simulate wave evolution for three types of bottom motions, i.e. impulsive, intermediate and slow bottom motion. The result displays a depression leading wave followed with subsequent wave crests. This negative wave explain the withdrawal of shoreline phenomena that we often observe preceding most tsunami events.

The organization of this paper is as follows. In Section 2 we first recall the Fourier and Laplace transform method for deriving the analytical transient wave. The solution is expressed in terms of complex integration that can be simplified using residue theorem. Here, we give detail derivation to yield an improper integral formula. In Section 3 we analyze the convergence of this integral and further we simulate transient waves as response to vertical bottom thrust using three different time scales. Finally in Section 4 we give concluding remarks.

2. ANALYTICAL SOLUTION OF LINEAR DISPERSION MODEL

For the sake of clarity, here we first recall the analytical formulation of the solution. Consider a layer of ideal fluid bounded below by a rigid bottom $z = -d(x, t)$ and above by free surface $z = \eta(x, t)$, as depicted in Figure 1. We note

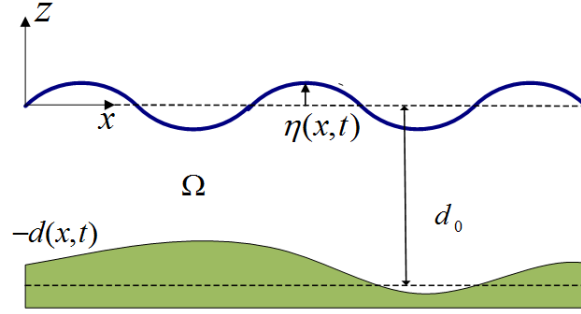


FIGURE 1. Schematic diagram of ideal fluid over a rigid bottom.

that to allow the bottom motion, we let it depends on time t . Under the assumption of irrotational flow, mass conservation for incompressible fluid read as

$$\phi_{xx} + \phi_{zz} = 0 \quad \text{in } \Omega, \quad (1)$$

where $\phi(x, z, t)$ is the velocity potential. Along the free surface, kinematic boundary condition and Bernoulli equation read as

$$\phi_z = \eta_t + \phi_x \eta_x \quad \text{on } z = \eta(x, t), \quad (2)$$

$$\phi_t + \frac{1}{2} |\nabla \phi|^2 + g\eta = 0 \quad \text{on } z = \eta(x, t). \quad (3)$$

Along the bottom, kinematic boundary condition that corresponds to impermeable bottom is

$$\phi_z = -d_t - \phi_x d_x \quad \text{on } z = -d(x, t). \quad (4)$$

In the derivation of analytical solution, we neglect the non linear term in (2)-(4). Furthermore the boundary conditions are considered on the undeformed free surface $z = 0$ and flat bottom $z = -d_0$. The linearized equations are explicitly

$$\phi_z = \eta_t \quad \text{on } z = 0, \quad (5)$$

$$\phi_t + g\eta = 0 \quad \text{on } z = 0, \quad (6)$$

$$\phi_z = -d_t \quad \text{on } z = -d_0. \quad (7)$$

Equations (5) and (6) are usually combined to yield

$$\phi_{tt} + g\phi_z = 0 \quad \text{on } z = 0. \quad (8)$$

The analytical solution for these problem given by equations (1),(6), (7) and (8) will be obtained by using integral transforms. We recall the definition of Laplace transform and its inverse as follows

$$\bar{f}(s) = \int_0^\infty e^{-st} f(t) dt \quad \text{and} \quad f(t) = \frac{1}{2\pi i} \int_\Gamma e^{st} \bar{f}(s) ds,$$

where Γ is a vertical line to the right of all singularities of $\bar{f}(s)$ in the complex s plane. Fourier transform and its inverse are defined as

$$\tilde{f}(k) = \int_{-\infty}^\infty e^{-ikx} f(x) dx \quad \text{and} \quad f(x) = \frac{1}{2\pi} \int_{-\infty}^\infty e^{ikx} \tilde{f}(k) dk.$$

In paragraph below we describe the steps in getting the analytical solution. We apply the Fourier transform w.r.t x and the Laplace transform w.r.t. t for equations (1), (6), (7) and (8) to yield

$$\tilde{\tilde{\phi}}_{zz}(k, z, s) - k^2 \tilde{\tilde{\phi}}(k, z, s) = 0, \quad (9)$$

$$\tilde{\tilde{\eta}}(k, s) = -\frac{s}{g} \tilde{\tilde{\phi}}(k, 0, s), \quad (10)$$

$$\tilde{\tilde{\phi}}_z(k, -d_0, s) + s\tilde{d}(k, s) = 0, \quad (11)$$

$$\tilde{\tilde{\phi}}_z(k, 0, s) + \frac{s^2}{g} \tilde{\tilde{\phi}}(k, 0, s) = 0. \quad (12)$$

General solution of the ordinary differential equation (9) is

$$\tilde{\tilde{\phi}}(k, z, s) = A(k, s) \cosh kz + B(k, s) \sinh kz. \quad (13)$$

In order to satisfy boundary condition (12), the relation between coefficients A and B should follow

$$B(k, s) = -\frac{s^2}{gk} A(k, s). \quad (14)$$

In order to get the coefficients A and B , we substitute (13) and (14) into equation (11) yielding

$$A(k, s) = \frac{gs\tilde{d}(k, s)}{(s^2 + \omega^2) \cosh kd_0}. \quad (15)$$

Solution (13) with coefficients (14) and (15) yield the following velocity potential in transformed variables

$$\tilde{\phi}(k, z, s) = \frac{gs\tilde{d}(k, s)}{(s^2 + \omega^2) \cosh kd_0} [\cosh kz - \frac{s^2}{gk} \sinh kz]. \quad (16)$$

Surface elevation in transformed variables can be obtained from (16) and (10) as follows

$$\tilde{\eta}(k, s) = \frac{-s^2\tilde{d}(k, s)}{(s^2 + \omega^2) \cosh kd_0}. \quad (17)$$

Next, applying the inverse Laplace and Fourier transforms of (17) to yield

$$\eta(x, t) = \frac{-1}{2\pi} \int_{-\infty}^{\infty} \left[\frac{1}{2\pi i} \int_{\Gamma} \frac{s^2 e^{-ikx} e^{st} \tilde{d}(k, s)}{(s^2 + \omega^2) \cosh kd_0} ds \right] dk. \quad (18)$$

In order to calculate (18) we must prescribe a specific bottom motion. For that purpose, we focus on a bed displacement governed by

$$d(x, t) = d_0 - \zeta_0(1 - e^{-\alpha t})H(b^2 - x^2), \quad (19)$$

where H is the heaviside step function. This motion is illustrated in Figure 2, in which the middle section of the bed $-b < x < b$ is moved downward to a maximum displacement ζ_0 in an asymptotic manner, prescribed by $(1 - e^{-\alpha t})$, in which α is related to a characteristic time t_c . A characteristic time t_c is the time required for the bottom to reach $\frac{2}{3}$ of the maximum displacement. This type of bottom motion is used in Hammack experimental works, as recorded in [1] and it is called the exponential bottom motion. The choice of the ratio $\frac{\zeta}{\zeta_0} = \frac{2}{3}$ is also used in [1].

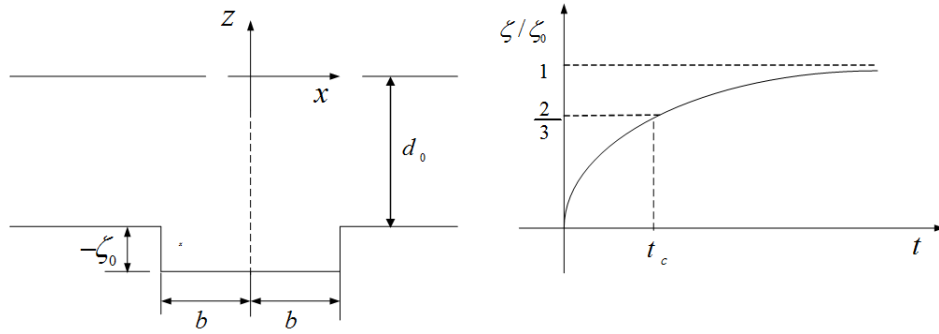


FIGURE 2. Bed displacement model.

The bed displacement (19) is then applied to (18) by expressing the Fourier and Laplace transforms of it

$$\tilde{d}(k, s) = -2\zeta_0 \frac{\sin kb}{k} \left[\frac{\alpha}{s(s + \alpha)} \right], \quad (20)$$

so that the surface elevation becomes

$$\eta(x, t) = \frac{-1}{2\pi} \int_{-\infty}^{\infty} \frac{-2\zeta_0 \sin kb}{k \cosh kd_0} e^{-ikx} \left[\frac{1}{2\pi i} \int_{\Gamma} \frac{\alpha s e^{st}}{(s^2 + \omega^2)(s + \alpha)} ds \right] dk. \quad (21)$$

In the next paragraph we give step by step of calculating the complex integral (21) using the residue theorem (see for reference [3] and [8]). We denote the integrand as $f(s) = \frac{\alpha s}{(s^2 + \omega^2)(s + \alpha)}$ which has three poles at $s = \pm i\omega$ and $s = -\alpha$, see the illustration in Figure 3.

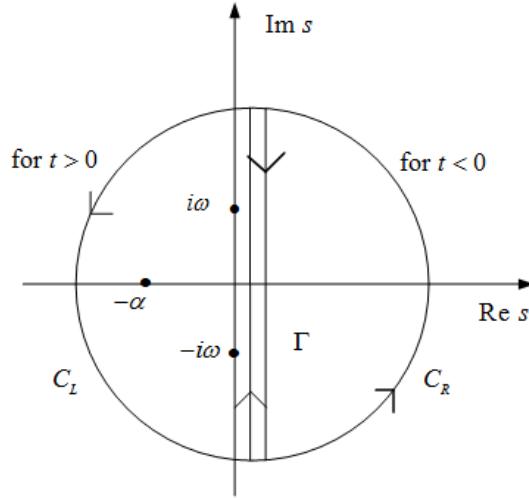


FIGURE 3. Contours for the inverse Laplace transform.

Let C_R and C_L are semi circles on the right and left half respectively, of the complex s plane. And since $C_R = \Gamma \cup C_1$ and $C_L = \Gamma \cup C_2$ as depicted in Figure 3, the complex integral in (21) can be evaluated as

$$\int_{\Gamma} f(s) e^{st} ds = \begin{cases} -\oint_{C_R} f(s) e^{st} ds + \oint_{C_1} f(s) e^{st} ds, & \text{for } t < 0, \\ \oint_{C_L} f(s) e^{st} ds - \oint_{C_2} f(s) e^{st} ds, & \text{for } t > 0. \end{cases} \quad (22)$$

Note that by Jordan's Lemma $\int_{C_1} f(s) e^{st} ds$ and $\int_{C_2} f(s) e^{st} ds$ are zero since

$$\lim_{s \rightarrow \infty} |f(s)| = 0.$$

Furthermore, with Cauchy's residue theorem, the integral (22) becomes

$$\int_{\Gamma} f(s) e^{st} ds = \begin{cases} 0, & \text{for } t < 0, \\ 2\pi i \sum_{j=1}^3 \operatorname{Res}_{s=s_{0j}} f(s) e^{st}, & \text{for } t > 0, \end{cases}$$

with

$$\sum_{j=1}^3 \operatorname{Res}_{s=s_{0j}} f(s)e^{st} = \operatorname{Res}_{s=i\omega} f(s)e^{st} + \operatorname{Res}_{s=-i\omega} f(s)e^{st} + \operatorname{Res}_{s=-\alpha} f(s)e^{st}.$$

The residue formula for a rational function at the simple pole $i\omega$, $-i\omega$, and $-\alpha$ will be used. Hence $\int_{\Gamma} f(s)e^{st} ds$

$$\begin{aligned} &= 2\pi i \left(\frac{\alpha s e^{st}}{K} \Big|_{s=i\omega} + \frac{\alpha s e^{st}}{K} \Big|_{s=-i\omega} + \frac{\alpha s e^{st}}{K} \Big|_{s=-\alpha} \right), \text{ where } K = 2s(s + \alpha) + s^2 + \omega^2 \\ &= 2\pi i \left(\frac{\alpha e^{i\omega t}}{2(i\omega + \alpha)} + \frac{\alpha e^{-i\omega t}}{2(-i\omega + \alpha)} + \frac{-\alpha^2 e^{-\alpha t}}{\alpha^2 + \omega^2} \right) \\ &= 2\pi i \left(\frac{\alpha^2 \cos \omega t + \alpha \omega \sin \omega t - \alpha^2 e^{-\alpha t}}{\alpha^2 + \omega^2} \right). \end{aligned}$$

Finally, the inverse transform reads

$$\frac{1}{2\pi i} \int_{\Gamma} \frac{\alpha s e^{st}}{(s^2 + \omega^2)(s + \alpha)} ds = \frac{\alpha^2}{\alpha^2 + \omega^2} \left(\cos \omega t + \frac{\omega}{\alpha} \sin \omega t - e^{-\alpha t} \right), \text{ for } t > 0. \quad (23)$$

Substituting (23) into (21) yield

$$\eta(x, t) = \frac{-\zeta_0}{\pi} \int_{-\infty}^{\infty} \frac{\sin kb}{k \cosh kd_0} e^{-ikx} \left[\frac{\alpha^2}{\alpha^2 + \omega^2} \right] \left(e^{-\alpha t} - \cos \omega t - \frac{\omega}{\alpha} \sin \omega t \right) dk. \quad (24)$$

Take only the real part of the resulting integral, and note that ω follows from the dispersion relation $\omega(k) = \sqrt{gk \tanh kd_0}$, moreover $\frac{1}{\omega^2} \cos \omega t$ and $\frac{1}{\omega} \sin \omega t$ are even functions in k , hence we have the water surface displacement resulting from the exponential bed motion as

$$\eta(x, t) = \frac{-2\zeta_0}{\pi} \int_0^{\infty} \frac{\cos kx \sin kb}{k \cosh kd_0} \left[\frac{\alpha^2}{\alpha^2 + \omega^2} \right] \left[e^{-\alpha t} - \cos \omega t - \frac{\omega}{\alpha} \sin \omega t \right] dk. \quad (25)$$

This is the explicit formula depending on the parameter of exponential bottom motion b , α , d_0 and ζ_0 .

3. NUMERICAL INTEGRATION RESULT

In this section, we use the analytical solution (25) to simulate free surface elevation as response of the downward bottom thrust. First, we note that analytical evaluation for the integral would be complicated, and it will be approximated by numerical integration. Let I denotes the integrand of (25). Integral in (25) is improper for two reasons, the infinite upper boundary, and singularity at $k = 0$. To handle this improper integral, we rewrite the integral (25) as $\int_0^{\infty} I dk = \int_0^1 I dk + \int_1^{\infty} I dk$. We analyze each of these integrals and we choose the integral which represents the analytical solution (25) for numerical integration.

We check the convergence of $\int_1^\infty Idk$ by first observe the following

$$\begin{aligned} |I| &\leq \frac{|\cos kx||\sin kb|}{|k \cosh kd_0|} \left| \frac{\alpha^2}{\alpha^2 + \omega^2} \right| \left(|e^{-\alpha t}| + |\cos \omega t| + \left| \frac{\omega}{\alpha} \sin \omega t \right| \right) \\ &\leq \frac{1}{|k \cosh kd_0|} \left| \frac{\alpha^2}{\omega^2} \right| \left(|e^{-\alpha t}| + 1 + \left| \frac{\omega}{\alpha} \right| \right) \\ &\leq \left| \frac{\alpha^2}{gk^2 \sinh kd_0} \right| (|e^{-\alpha t}| + 1) + \left| \frac{\alpha}{k^{\frac{3}{2}} \sqrt{(g \sinh 2kd_0)/2}} \right| \end{aligned}$$

Using comparison test with the convergent integral $\int_1^\infty \frac{1}{k^{\frac{3}{2}}} dk$, we get the convergence of $\int_1^\infty \left| \frac{\alpha}{k^{\frac{3}{2}} \sqrt{(g \sinh 2kd_0)/2}} \right| dk$. Similar argument with the convergent integral $\int_1^\infty \frac{1}{k^2} dk$, we can argue that $\int_1^\infty \left| \frac{\alpha^2}{gk^2 \sinh kd_0} \right| (|e^{-\alpha t}| + 1) dk$ is also converge. The convergence of these components, yield the convergence of $\int_1^\infty Idk$.

The convergence and value of this integral is also check by calculating its value at a chosen position and time, $x = 61$ and $t = 2$, by means of trapezoidal numerical integration. Their values are displayed in Table 1. For example, the first data in Table 1 show $\int_1^{1.1} I(61, 2) dk = 1.1034 \cdot 10^{-6}$. It is shown in the right column, the value of $\int_1^\infty Idk$ is close to zero, with accuracy up to 10^{-6} which suggest that $\int_1^\infty Idk \approx 0$. We also check the value of this integral at others position and time.

Next, we focus on applying numerical integration for calculating $\int_0^1 Idk$. To handle the singularity of the integrand at $k = 0$ we first observe its limit value. By L'Hopital's rule $\lim_{k \rightarrow 0} I = b(e^{-\alpha t} - 1)$. For numerical integration, we apply the trapezoidal rule whereas the value of I at $k = 0$ is taken to be its limit value $b(e^{-\alpha t} - 1)$. In Table 2 we calculate $\int_a^b I(61, 2) dk$ for various integration interval in $[0, 1]$. For example, the first data in Table 2 show $\int_0^{0.1} I(61, 2) dk = -0.49005716$. Finally, after analyze the integral $\int_1^\infty Idk$ and $\int_0^1 Idk$, we conclude that integral in (25) can be approximated as $\int_0^\infty Idk \approx \int_0^1 Idk$.

TABLE 1. Calculating $\int_a^b I(61, 2) dk$ using trapezoidal rule with $\Delta x = 0.0001$, for $[a, b] \subset [1, 1000]$

$[a, b]$	Value
$[1, 1.1]$	$1.1034E - 06$
$[1, 1.2]$	$1.0715E - 06$
$[1, 1.5]$	$1.0750E - 06$
$[1, 2]$	$1.0730E - 06$
$[1, 10]$	$1.0730E - 06$
$[1, 100]$	$1.0730E - 06$
$[1, 1000]$	$1.0730E - 06$

TABLE 2. Calculating $\int_a^b I(61, 2)dk$ using trapezoidal rule with $\Delta x = 0.0001$, for $[a, b] \subset [0, 1]$

$[a, b]$	Value
$[0, 0.1]$	-0.49005716
$[0, 0.2]$	-0.47260392
$[0, 0.5]$	-0.47393012
$[0, 0.6]$	-0.47383404
$[0, 0.7]$	-0.47393062
$[0, 0.8]$	-0.47391938
$[0, 1]$	-0.47391243

Next, we calculate $\eta(61, 2)$ as $\int_0^1 I(61, 2)dk$ using various length of Δx as shown in Table 3. We conclude that taking $\Delta x = 0.0001$ is enough to get the correct result up to 10^{-5} .

TABLE 3. Calculating $\eta(61, 2) \approx \int_0^1 I(61, 2)dk$ using various length of Δx .

Δx	N	$\int_0^1 I(61, 2)dk$
0.1	10	-2.90747104
0.05	20	-1.41682363
0.02	50	-0.47391186
0.01	100	-0.47391234
0.001	1000	-0.47391243
0.0001	10000	-0.47391243

In later calculations, we use integration interval $[0, 1]$ with $\Delta x = 0.001$. Here, we simulate surface wave generated by specific bottom motion. We use the exponential bottom motion (19) with parameter $d_0 = 5$, $b = 61$ and $\zeta_0 = 1$. We choose the characteristic time $t_c = 0.08$. The explicit analytical formula (25) is defined for $t > 0$ and $x \in R$ presents the free surface elevation as an even function in x . In Figure 4 we plot the snapshots of surface elevation on the half interval $x > 0$ for subsequent time $t = 0.04$, $t = 0.1$, $t = 0.25$, $t = 0.5$, $t = 1$ and $t = 1.2$ near the characteristic time. The result shows the surface wave will increase correspond with bottom movement and after the characteristic time, it will form a negative leading wave.

In Figure 5 we plot the snapshots of surface elevation on the half interval $x > 0$ for subsequent time $t = 5$, $t = 10$, $t = 15$ and $t = 20$ when bottom motion is complete. The result show that the free surface responds in terms of a negative leading wave, followed with dispersive train of waves. Moreover, we observe that the maximum amplitude of the surface is 0.6 which is approximately half of the bottom amplitude ζ_0 . And this follow the analytical prediction as described in [2]. Moreover, the leading wave moves according to the linear velocity $\sqrt{gd_0}$.

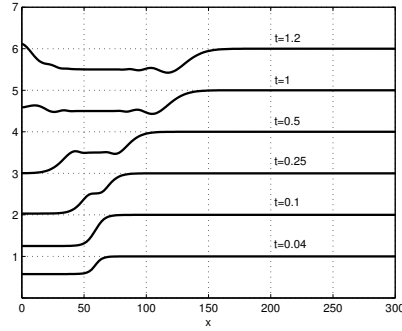


FIGURE 4. Snapshots of analytical surface elevation as response of a downward bottom thrust, at the early stage of generation process.

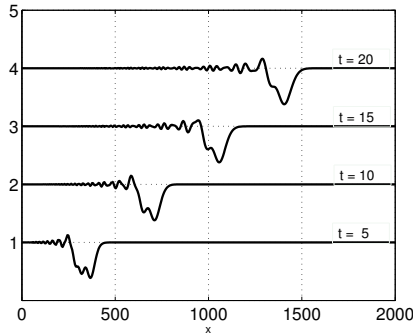


FIGURE 5. Snapshots of analytical surface elevation after generation being completed. The developed wave experiences only slight deformation during its evolution over flat bottom

Further, we investigate the effect of characteristic time t_c . The same bottom motion and parameter are used, with difference only on the choice of the characteristic time t_c . We use some different characteristic times, $t_c = 0.08$, $t_c = 0.8$, and $t_c = 4.0$, that represent three different kind of bottom motions: impulsive, intermediate and slow motion, respectively. The results at $t = 10$ are plotted together in Figure 6. We observe that slower bottom motion yield smaller wave amplitude. Numerical results also suggest that for a fixed downward bottom thrust, the impulsive motion is the one that has the most dangerous effect. Moreover, only for impulsive type of bottom motion, the surface wave has similar form with bottom profile.

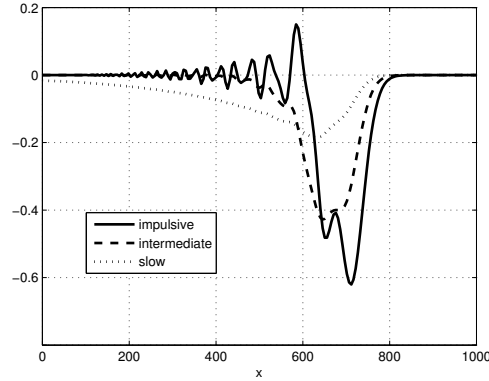


FIGURE 6. Surface wave profiles at time $t = 10$ for impulsive, intermediate and slow bottom motions.

4. CONCLUDING REMARKS

We have presented the explicit analytical formulation of surface elevation as response of bottom motion in terms of improper integral. Analyzing the convergence of this integral, and applying trapezoidal numerical integration have given us the development of free surface wave. Simulation of the downward bottom thrust yields a negative leading wave, which is followed with a dispersive wave train. The explicit analytical formula allowing us to simulate transient wave for different type of motion. Among three types of motions: impulsive, intermediate, and slow, the impulsive motion yield the maximum amplitude of leading wave. In the case of impulsive bottom motion, the amplitude and velocity of our leading wave confirm the analytical formula. It should be noted that once the wave has been generated, our description does not considered topographic change during the evolution. So, for predicting wave characteristic at the beach, further investigations are still needed. Also, our analysis here was only for a downward bottom motion.

Acknowledgement. Financial support from Riset Desentralisasi ITB 2015, 311n/I1.C01/PL/2015 and 310c/I1.C01/PL/2015 are greatly acknowledged. Partial support from the Ensemble Estimation of Flood Risk in a Changing Climate project funded by The British Council through their Global Innovation Initiative are also acknowledged.

REFERENCES

- [1] Hammack, J.L., *Tsunamis-a model of their generation and propagation*, W.M. Keck Lab. Hydraul. and Water Res., Calif. Inst. Tech. Rep. KH-R-28, 1972.

- [2] Hammack, J.L., "A note on tsunami: their generation and propagation in an ocean of uniform depth", *J. Fluid Mech.*, **4** (1973), 769-799.
- [3] Mei, C.C., and Yue, D.P., *Advanced series on Ocean Engineering*, World Scientific, 2005.
- [4] Dutykh, D., and Dias, F. *Tsunami and non linear waves*, Springer, 2007.
- [5] Kervella, Y., Dutykh, D., and Dias, F., "Comparison between three dimensional linear and non linear tsunami generation models", *Theor. Fluid Dyn.*, **21** (2007), pp. 245-269.
- [6] Fuhrman, D.R., and Madsen, P.A., "Tsunami generation, propagation, and run-up with a high-order Boussinesq model", *Coastal Engineering*, **56** (2009), pp. 747-758.
- [7] Lynett, P., and Liu, P.L.F., "A numerical study of submarine-landslide-generated waves and run-up", *Proc. R. Soc. Lond. A*, **458** (2002), pp. 2885-2910.
- [8] Kreyszig, E., *Advanced Engineering Mathematics*, John Wiley and Sons, 1999.