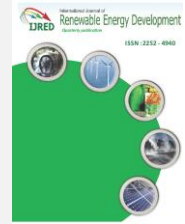




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Variations of Wave Energy Power in Shoaling Zone of Benin Coastal Zone

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ABSTRACT: Today, we observe at the population level, that the improvement in comfort is accompanied by an increase in the electrical energy required. The predicted exhaustion of fossil energy resources maintains some speculation. Their unequal geographical distribution justifies the energy dependence of Benin overlooked from outside. So it is urgent to explore the various sources of renewable energy available to Benin. In this work, using measurements made by the Millennium Challenge Account (MCA-Benin) as part of the extension of the port of Cotonou, with Boussinesq equations (Peregrine) and Stokes waves dispersion relation, we characterized the variations of various swell parameters (height, wavelength, velocities) in the shoaling zone on the study site and proceeded to estimate variations in wave energy power from deep waters to the bathymetric breaking point. Finally, the zone with high energy power (where the conversion of this energy into electrical energy would be profitable) of these waves is highlighted on the site, the local water depth at the point of breaking waves is evaluated and results obtained allowed to justify the very energetic character take by these swells on this coast when they are close to the beach.

Keywords: wave energy power; weakly nonlinear waves; Benin coastal zone; wave shoaling; nonlinear Boussinesq equations

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1. Introduction

The description of the ocean, fluid medium, heterogeneous, turbulent, headquarters of numerous interactions at different scales of time and space, is not accessible without a dual approach of geographical exploration and scientific understanding. It is even more complex when the swell is influenced by the seabed that causes the amplification of its height and the contraction of its wavelength in the coastal areas (Babarit, 2009; Battjes, 2004). These variations of the different swell parameters imply variations in its energetic power depending on the nature and the slope of the seabed.

Benin energy situation is characterized by low access to electricity and almost total dependence on outside electrical energy rates. Coastal nation of Gulf of Guinea, it benefits to the south the access to the ocean

for a distance of 125 km approximately (Houékpohéha et al., 2014). On this side, the almost permanent propagation of regular waves is observed. To ensure energy independence, the exploration of all available sources of renewable energy is essential. An advantageous solution to fight against coastal erosion in Benin and the sub-region is the transformation of wave energy into electrical energy (Kounouhéwa et al., 2014). To this end, highly accurate and consistent characterization of this energy resource is crucial in order to identify areas of high energy output.

In the course of its propagation approaching the coast, swell undergoes the modifications related to the depth of decreasing and coastal morphology (Barthe et al., 1989). When the wave approaches the coast, it undergoes deformations according to schematic evolution below (Figure 1) (Astier, 2014): at the same

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time she stiffened, water depth decreases (lifted zone or shoaling zone) to the breaking point associated with intense turbulent motions (breaking zone), then the waves reorganize themselves in the form of propagating hydraulic jump (internal surf zone), which reach the beach by generating run-up (Swash zone) (Bonneton, 2003 and Bird, 1985). In previous work, Bonneton (2003), Panchang (1991) and others have shown that the basic parameters that characterize the swells are their wavelength (L), their period (T), the vertical elevation of the sea surface (η) or peak to valley height (H), the phase velocity (C_ϕ) and the group velocity (C_g) of wave (Bonneton et al., 2001) and (Panchang et al., 1991).

The coastal zone where the swell feels the perturbation effect of the seabed is divided into two areas: the intermediate zone (which is offered the shoaling wave) and the surf and swash zones (Kounouhéwa et al., 2014). After identifying, from the measurement data made, the equations that describe the swells in the shoaling zone of the coastal zone of Benin, the objective in this work is to evaluate the variations of parameters that characterize swells, to estimate changes in their energy power in this area and to predict wave height and local water depth at bathymetric breaking point of these swell is performed in order to identify areas of high energy power where the installation of a converter of this energy into electrical energy could be profitable.

2. Methodology

2.1. Presentation of the study site and data used:

Benin is a country in the Gulf of Guinea between the parallel $6^{\circ}15'$ and $12^{\circ}30'$ north latitude one hand and meridians 1° and $3^{\circ}40'$ east longitude other hand. It has a coastal area with a length of 125 kilometers from Hilla-Condji in the west to Kraké in the east. Benin's coast is more or less linear and interrupted at two locations, namely: the Mouth of Roy and the mouth of the Cotonou channel.



Figure 1. Geographic situation of Benin coastal zone

Benin coastal zone is between $6^{\circ}15'$ and $6^{\circ}23'$ to north latitude. It is part of the coastal sedimentary basin of which the oldest date back to the Cretaceous period lands (the end of the secondary era). The climate is sub-equatorial with two dry seasons and two rainy seasons alternate (Houékphéha et al., 2014)

This coastal zone, belonging to Gulf of Guinea, receives regular waves of low amplitude. Benin bathymetry and wave power measurements in six months time step of five minutes as part of the extension of the port of Cotonou by the Millennium Challenge Account (MCA-Benin) helped define stable average data on the site of the port of Cotonou.

- In deep waters, the average wave height is $H_o = 2a_o = 0.8 \text{ m}$ and its stable average period $T = 12 \text{ s}$ at a position where the local water depth is approximately $d_o = 130 \text{ m}$.
- The gravity on the Benin coast is approximately $g = 9.79 \text{ m.s}^{-2}$ and wavelength in deep waters is approximately $L_o = 200 \text{ m}$ (Dégbé et al., 2010)
- On this site, the period of the waves varies between 8s and 18s and its stable average value is 12s. This parameter is the only one that remains practically constant during a given wave shoaling. ($T \approx T_o = \text{cst}$)
- The bathymetric map (below) obtained near the IRHOB (Institute of Halieutic and Oceanographic Research of Benin) shows the evolution of local water depth in Benin coastal zone and allowed us to predict the average slope of the seabed

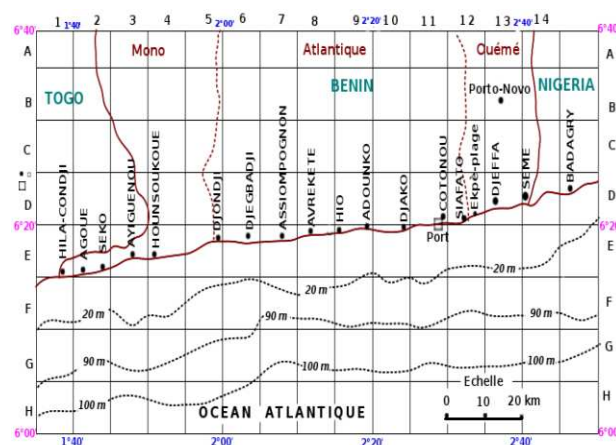


Figure 2. Bathymetric map of Benin seabed

- The seabed in the coastal zone is almost (less irregular) flat and inclined. It is a gently sloping seabed $p = \tan \beta$ such as $0.001 < \tan \beta \leq 0.1 \Rightarrow \tan \beta \approx \beta$. The average of the slope in the shoaling zone is $\beta_m \approx \frac{100}{20000} = 0.05$ (Kounouhéwa et al., 2014b).
- A series of measurements made on the site shows that the local water depth at the breaking point

varies between 0.8 m and 2.7 m while height at this point is between 1 m and 2.6 m.

2.2. Variations of wave parameters from deep waters to breaking point

For waves whose wavelength is L , three propagation zones are defined: deep waters when $d \geq \frac{L}{2}$ and shoaling zone if $\frac{L}{25} \leq d \leq d_b$ (Kounouhéwa et al., 2014)

The dispersion relation of Airy or Stokes waves is (Kounouhéwa et al., 2014a; Fenton, 1990)

$$\omega^2 = \frac{4\pi^2}{T^2} = gk \tanh(kd) \quad (1)$$

Expressions reflecting changes in phase velocity C_ϕ and group velocity C_g are: (Kounouhéwa et al. 2014a; Catalan and al. 2008).

$$\begin{cases} C_\phi = \frac{\omega}{k} = \sqrt{\frac{g}{k} \tanh(kd)} \\ C_g = \frac{1}{2} \left(1 + \frac{2kd}{\sinh(2kd)} \right) \sqrt{\frac{g}{k} \tanh(kd)} \end{cases} \quad (2)$$

In deep waters ($d \geq \frac{L_o}{2} \Rightarrow \mu \geq \frac{1}{2}$), $\tanh(kd) \approx 1$ then all parameters are constant in this area of propagation and expression of the wavelength, phase and group velocities are:

$$L_o = \frac{gT^2}{2\pi} \text{ and } C_{\phi_o} = 2C_{g_o} = \frac{gT}{2\pi} \quad (3)$$

In the Gulf of Guinea at Cotonou, the waves are Stokes swells ($\eta \ll d$), weakly nonlinear ($\varepsilon \ll 1$) and low-dispersive ($\mu^2 < 1$).

In Benin shoaling zone ($d_b \leq d \leq \frac{L_o}{2} \Rightarrow \mu_b \leq \mu \leq \frac{1}{2}$), the nonlinearity parameter $\varepsilon = \frac{H}{2d}$ increase, seabed is nearly flat and Stokes waves become nonlinear. As the local water deep is variable, the wave propagation in 2D is modeled by the Boussinesq equations (Peregrine): (Chazel, 2007) and (Bellec et al., 2013)

$$\begin{cases} \frac{\partial \eta}{\partial t} + \frac{\partial (du)}{\partial x} = 0 \\ \frac{\partial u}{\partial t} + g \frac{\partial \eta}{\partial x} = \mu^2 \left[\frac{d}{2} \frac{\partial^2}{\partial x^2} \left(d \frac{\partial u}{\partial t} \right) - \frac{d^2}{6} \frac{\partial^3 u}{\partial x^2 \partial t} \right] + \mathcal{O}(\mu^4) \end{cases}$$

$$\text{with } \begin{cases} h = \eta + d \\ u = u(x, t) \\ \eta = \eta(x, t) \\ d = d(x) \end{cases} \quad (4)$$

Vertical elevation $\eta(x, t)$ of ocean free surface is given by

$$\eta(x, t) = \frac{H(x)}{2} e^{i(kx - \omega t + \phi)} \quad (5)$$

Boundary conditions at the border of deep waters and the beginning of shoaling zone where $\mu = \frac{1}{2}$ or $d = \frac{L_o}{2}$, are:

$$H\left(\mu = \frac{1}{2}\right) \approx H_o \text{ and } L\left(\mu = \frac{1}{2}\right) \approx L_o \quad (6)$$

The equation of tangent line to the seabed is $z = -d = x \tan \beta$ (Tsai et al., 2002; Zou et al., 2003) with $d = \mu L_o$ (Senechal, 2003). On the study site $\tan \beta \approx \beta$, thus:

$$d = -\beta x = \mu L_o \quad (7)$$

Using the relationship (5), (6) and (7) to solve Boussinesq (Peregrine) equations, we obtain in shoaling zone:

$$\begin{cases} H = H_o \left(\frac{1}{2\mu} \right)^{1/4} = H_o \left(\frac{2d}{L_o} \right)^{-1/4} \\ L = L_o (2\mu)^{1/2} = L_o \left(\frac{2d}{L_o} \right)^{1/2} \end{cases} \quad (8)$$

According to this theory, which takes into account the variability of the bathymetry, the different coefficients $\delta_1 = \frac{C_\phi}{C_{\phi_o}}$, $\delta_2 = \frac{C_g}{C_{g_o}}$, $\delta = \frac{L}{L_o}$ and $K_s = \frac{H}{H_o}$ respectively show the changes of phase velocity, group velocity, wavelength and height of the swell from deep waters to the bathymetric breaking point.

2.3. Variations in swell energy power:

The average of the total energy E per unit length of a swell of which the height is H is expressed: (Ardakani et al., 2010)

$$\langle E \rangle = \frac{1}{8} \rho g H^2 \quad (9)$$

The energy power P per unit length of a wave is the product of its average total energy and its group velocity (Ardakani et al. 2010; Babarit et al. 2009)

$$P = \langle E \rangle \cdot C_g \quad (10)$$

In deep waters, we have:

$$P = \langle E \rangle_o \cdot C_{g_o} = \frac{1}{8} \rho g H_o^2 C_{g_o} = \frac{1}{32\pi} \rho g^2 H_o^2 T \quad (11)$$

Thus $\chi = \frac{P}{P_o}$ is the coefficient which reflects the variations in energy power from deep waters to the breaking point.

2.4 Wave height and local water deep at breaking point

Kaminsky and Kraus, (1993) and Sabatier, (2001) after conducting a comparative study based on an analysis of previous studies (409 cases) covering a wide range of wave cambers and beach slopes, propose to use for $\tan\beta \leq 0,1$ the breaking index below that takes into account Irraben number $\left(\xi = \tan\beta \sqrt{\frac{L_o}{H_o}}\right)$ (Tsai and al., 2002)

$$\gamma = \frac{H_b}{d_b} = 1.2 \left(\tan\beta \sqrt{\frac{L_o}{H_o}} \right)^{0,27} \quad (12)$$

On the study site, $0.001 < \tan\beta \leq 0.1$, this criterion is selected to determine the breaking point of the waves and we obtain.

$$\begin{cases} d_b = \frac{H_o^{0,908} L_o^{0,092}}{1,33 (\tan\beta)^{0,216}} \\ H_b = 0,903 H_o \left(\frac{L_o}{H_o} \right)^{0,227} (\tan\beta)^{0,054} \end{cases} \quad (13)$$

3. Results and Discussion

The curves in Figure 3 are the coefficients of variation of the wavelength, of height, of the phase and group velocities and energy power of swell in the Gulf of Guinea in Benin.

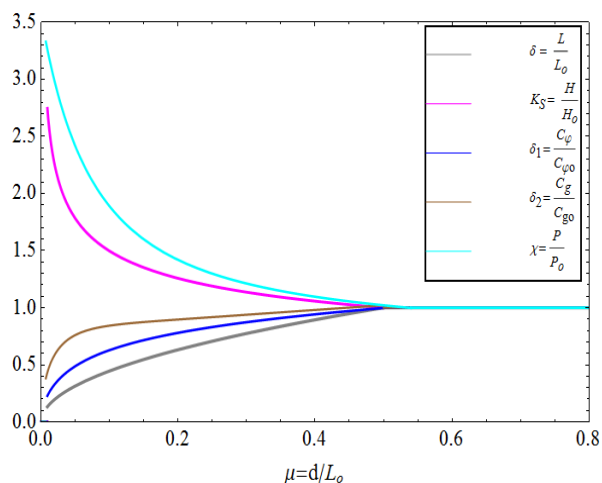


Figure 3. Coefficients of variations of wave characteristics parameters in the Gulf of Guinea in Benin.

All curves in Figure 3 show that all wave parameters in deep waters, apart from any obstacle, are constant.

The curve $K_s = \frac{H}{H_o}$ represents the variation of the height in deep waters and in the shoaling zone. It is constant in deep waters ($K_s \approx 1$) and believes strictly in the shoaling zone ($1 \leq K_s \leq 2.8$) and this amplification stops at the breaking point. The peak to valley wave's height, constant in deep waters is strictly increasing in the shoaling zone. This height, which believes strictly when the local water depth decreases, remains proportional to $(d^{-1/4})$: this result coincides with the results obtained for the shoaling of subharmonic gravity waves (Battjes et al. 2004).

Curves $\delta = \frac{L}{L_o}$, $\delta_1 = \frac{C_\phi}{C_{\phi_o}}$ and $\delta_2 = \frac{C_g}{C_{g_o}}$ decreases in shoaling zone when the local water depth decreases.

But the decrease was more pronounced in the case δ and low for δ_2 . In the coastal zone of the Gulf of Guinea in Cotonou, the wavelength of swell decreases when the local water depth decreases and remains proportional to the square root of the latter ($L \sim \sqrt{d}$).

The evolution of the curve δ_2 shows that the group velocity decreases strictly in the shoaling zone. This result justifies the fact that the speed of a system subject to energy dissipation cannot grow

Variations of δ_1 indicate that the phase velocity decreases as the waves move towards the shore. The phase and group velocities of swells in the shoaling zone in Cotonou decrease when the local water depth decreases. But in this area, the group velocity is always greater

The curve $\chi = \frac{P}{P_o}$ reflects variations in wave energy power. In deep waters, this curve shows that this quantity is constant. In shoaling zone, it is amplified when the local water depth decreases and this amplification stops at the breaking point. It shows that $1 \leq \chi \leq 3.5$: Energy power can be amplified 3.5 times before the breaking point.

The following diagram in Figure 4 shows the variations in the available energy power waves in deep waters and in the shoaling zone on the Benin coast.

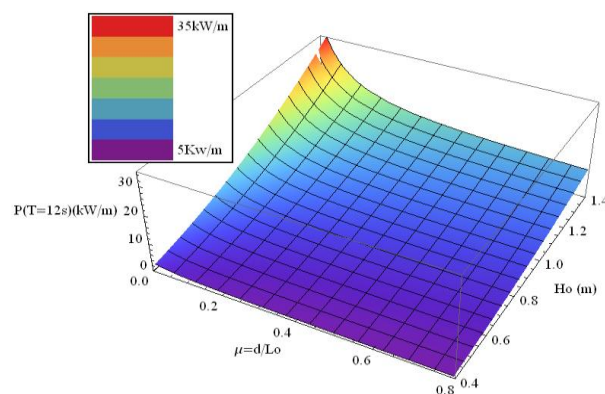


Figure 4. Variations in wave energy power in the coastal zone of Benin

The evolution of the diagram in Figure 4 shows that the energy available power of the waves in the coastal zone of Benin is amplified when the local water depth decreases to the breaking point. In deep waters, the power varies between. 5 kW/m and 12 kW/m and approximately. Under the effect of shoaling, the energy power and believes and oscillates between 10 kW/m and 35 kW/m : These swells therefore become more energetic when the seabed slope increases.

The curves in Figure 5 herein after, show the evolution of the local water depth d_b and the wave height H_b up to the bathymetric breaking point depending on the wave's period T one hand and depending on the height of the swells off the other.

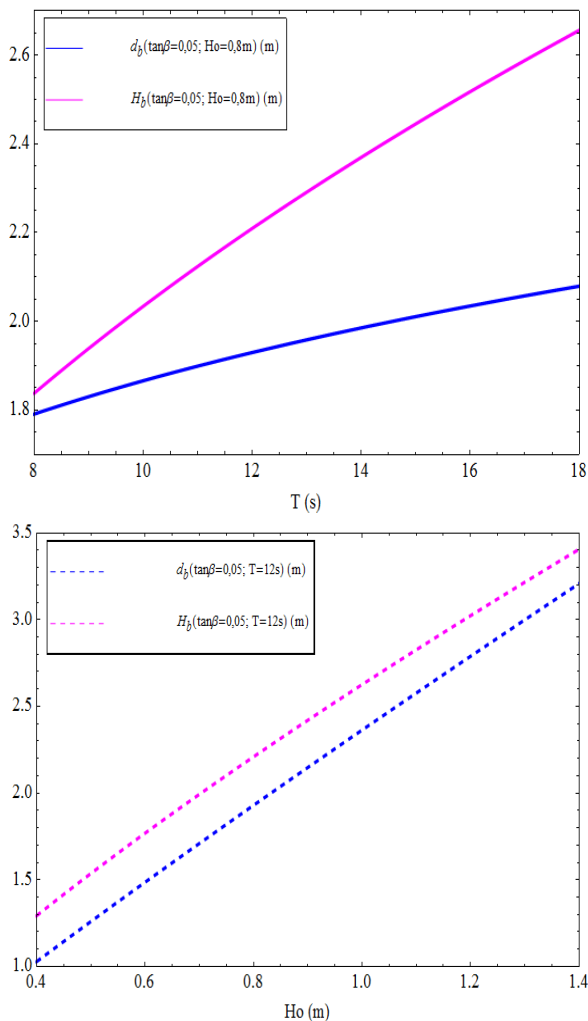


Figure 5. Variations of local water depth d_b and wave height H_b at the breaking point in 2D according T or H_o

As to diagrams in Figure 6, they show respectively, for a seabed slope $\tan\beta = 0.05$, changes in the local water depth d_b and those of the height H_b of waves at the bathymetric breaking point.

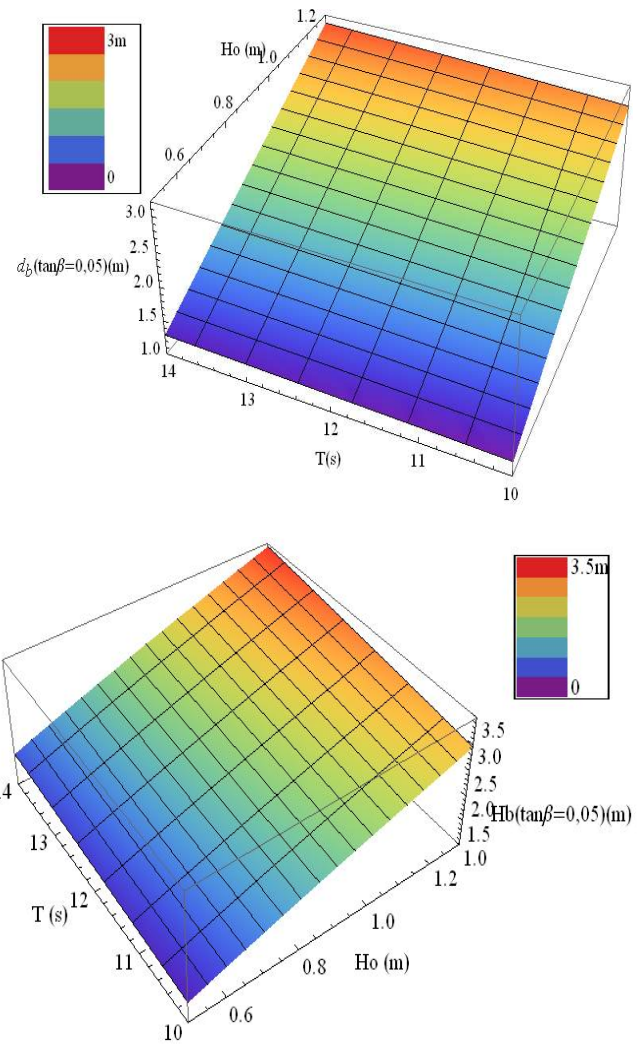


Figure 6. Variations of local water depth d_b and wave height H_b at the breaking point in 3D according T and H_o

The curves in Figure 5 and Figure 6 diagrams obtained from the results of the theory of Boussinesq and breaking criterion Kaminsky and Krauss (1993) show that the local water depth and height swell to the breaking point varies depending on the slope of the seabed, period and wave height in deep waters. Greater the slope, the less great is the local water depth at the point breaking and greater is the height of the waves at this point. These results show that, in the coastal zone of the Gulf of Guinea in Cotonou on bathymetric wave breaking occurs at a position where the local water depth varies between 1 m and 3.2 m ($1 \text{ m} \leq d_b \leq 3.2 \text{ m}$) when $0.4 \text{ m} \leq H_o \leq 1.4 \text{ m}$ and $10 \text{ s} \leq T \leq 14 \text{ s}$: greater the slope, the greater the local water depth at breaking point decreases. Under the same conditions, the wave height at breaking point varies between 1.2 m and 2.8 m approximately ($1.2 \text{ m} \leq H_b \leq 2.8 \text{ m}$). These results are very similar measurements performed ($0.8 \text{ m} \leq d_b \leq 3 \text{ m}$ and $1 \text{ m} \leq H_b \leq 2.7 \text{ m}$) and justify the appropriateness of breaking criterion is chosen.

Figure 7 diagrams below; reflect variations in the vertical elevation of the free surface according to the local water depth $d = \mu L_o$ and slope $\tan\beta = 0,05$ of the seabed.

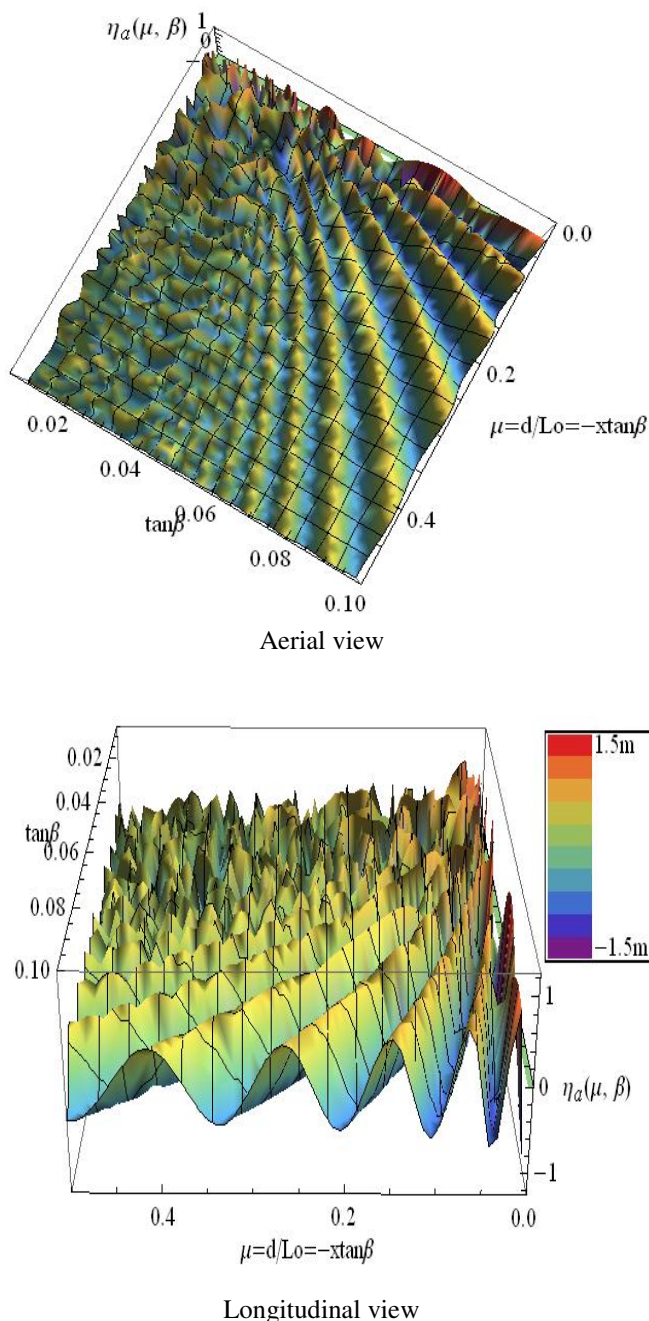


Figure 7. Vertical elevation of the sea surface in the shoaling zone in 3D according to the local water depth $\mu = d/L_o$ and the seabed slope β

Their variations confirm:

- The consistency of the various parameters in deep waters.

- Amplification of the peak to valley height of waves in shoaling zone disrupted the seabed.
- The decrease or contraction of the wavelength caused by the slope of the seabed.
- The fluctuating and randomness of the deformation of the surface of the ocean, which is observed on the coast of Benin, even in the absence of a local wind. These results indicate that the variability of slope of the seabed helps randomness of oscillations at the ocean surface in the coastal zone.

From this work, it emerges that in the coastal zone of Benin (in the Gulf of Guinea), the swells become very energetic when they move towards the coast. Their bathymetric breaking which comes very close to the coastline and coastal ground geology in Cotonou (non-cohesive sediments: sandy soil) are the major causes of coastal erosion which is observed along this site.

4. Conclusion

In the Gulf of Guinea at Cotonou, the waves are regular and have a constant height $H_o = 0.8\text{ m}$ and a mean period $T = 12\text{ s}$ in the deep waters. In the coastal zone, the perturbation effect of the seabed shoaling causes their up to breaking point. These waves become very energetic in this area; their height is amplified and is proportional to $(d^{-1/4})$. Under the effect of shoaling, the wave energy power is amplified when the local water depth decreases and can vary between 10 kW/m and 35 kW/m . At break point, the maximum height reached by the waves varies between $1.2\text{ m} \leq H_b \leq 2.3\text{ m}$. Their depth breaking occurs at a position where the local water depth d_b , which oscillates between $1\text{ m} \leq d_b \leq 3.2\text{ m}$ according to the value of the seabed slope, is very close to the coastline. This breaking is a brutal energy discharge that induces and increases the phenomenon of coastal erosion. Benin and therefore the sub-region have two advantages to be gained from converting this energy:

- Curbing the cyclic crisis of electrical energy in the sub-region;
- Mitigate Coastal Erosion in the Gulf of Guinea by the absorption of sudden energetic discharge of waves thanks to various conversion systems of marine energy.

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Nomenclature of parameters

$\eta(x, t)$: Vertical ocean surface elevation (m)
H	: Wave peak to valley height (amplitude) (m)
T	: Wave period (s)
L	: Wavelength (m)
C_ϕ	: Wave phase velocity (m/s)
C_g	: Wave group velocity (m/s)
d	: Local water deep (m)
$\tan \beta$: Seabed slope
$\mu = \frac{d}{L_0}$: Dispersion parameter of waves train
$\varepsilon = \frac{2H}{d}$: Nonlinearity parameter
P	: Wave energy power (W/m^2)
$\delta = \frac{L}{L_0}$: Coefficient of wavelength variations
$\delta_1 = \frac{C_\phi}{C_{\phi_0}}$: Coefficient of phase velocity variations
$\delta_2 = \frac{C_g}{C_{g_0}}$: Coefficient of group velocity variations
$K_s = \frac{H}{H_0}$: Coefficient of wave height variations (shoaling coefficient)
$\chi = \frac{P}{P_0}$: Coefficient of wave energy power variations
H_b	: Wave height at breaking point (m)
d_b	: Local water deep at breaking point (m)

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